

dSECS

Including Variometers in Geomagnetic Field Interpolation

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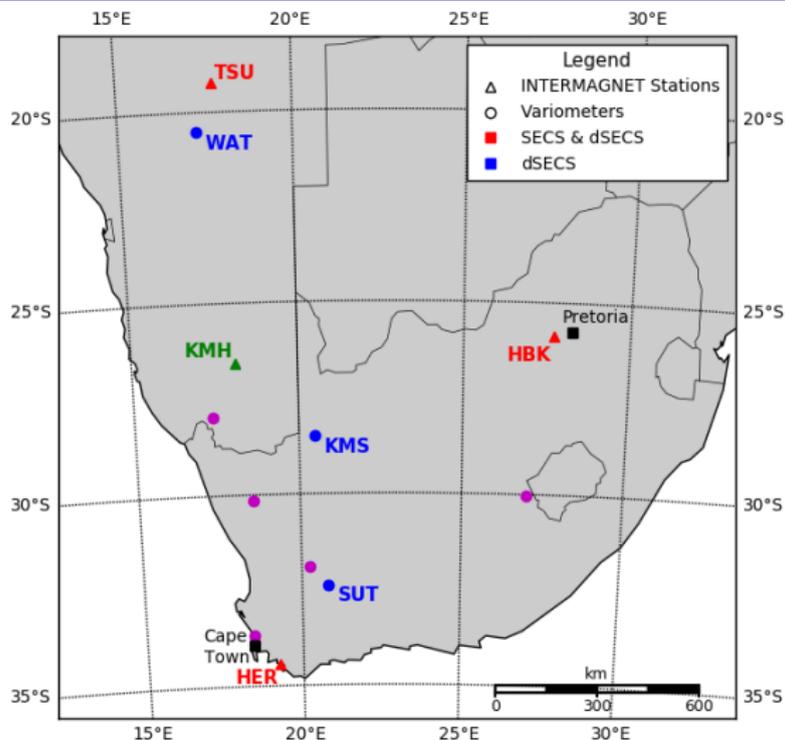
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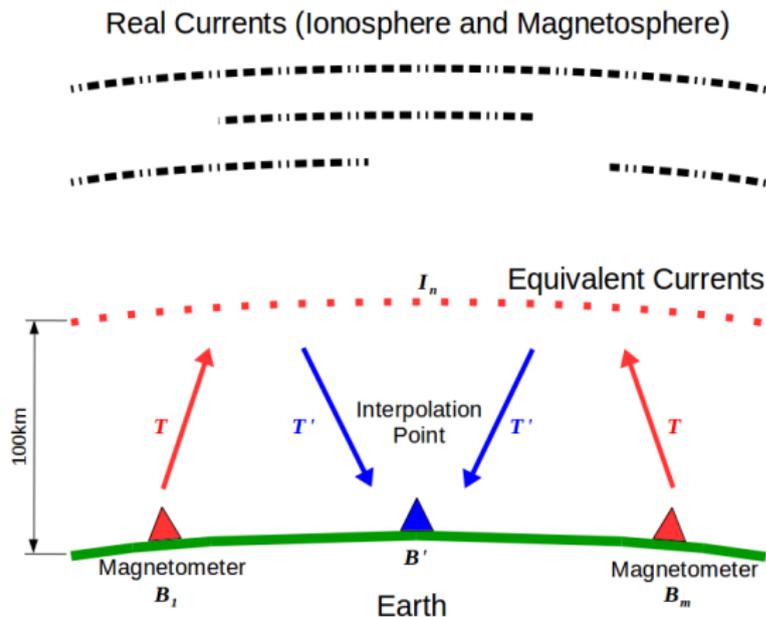


Variometers and Magnetometers



SECS - Spherical Elementary Current Systems

SECS is one of the most physically realistic magnetic field interpolation schemes - it tries to replicate both the internal (induction) and external (ionospheric and magnetospheric) currents in terms of equivalent currents.



SECS - Algorithm Flow

- 1 - Set up grid of elementary current poles (say n poles). These poles are collected into a vector I .
- 2 - Use measured magnetic field data (say m stations) to constrain elementary currents. Usually $m < n$. These stations are collected into a vector B .
- 3 - Set up transfer function matrix T , such that $B = T \cdot I$. The dimensions of T is $m \times n$
- 4 - Calculate components of I by using SVD to get T^{-1} , i.e. $I = T^{-1} \cdot B$.
- 5 - Use the elementary current grid to interpolate magnetic field B' to any other point of of interest, i.e. $B' = T' \cdot I$.

The internal component of the measured geomagnetic field needs to be subtracted!



SECS - Planar Approximation

Given a grid with n current elements and assuming m magnetometers, the following matrix equation can be set up for the calculation,

$$\begin{bmatrix} B_{x,y:1} \\ \vdots \\ B_{x,y:m} \end{bmatrix} = \begin{bmatrix} T_{x,y:11} & \cdots & T_{x,y:1n} \\ \vdots & \ddots & \vdots \\ T_{x,y:m1} & \cdots & T_{x,y:mn} \end{bmatrix} \begin{bmatrix} I_1 \\ \vdots \\ I_n \end{bmatrix}$$

where,

$$T_{x,y:ij} = \frac{\mu_0}{4\pi r} \left(1 - \frac{h}{\sqrt{r_{ij}^2 + h^2}} \right).$$

There are a number of factors affecting this planar approximation including an small angle assumption and the resolution of stations.



dSECS - Interpolating dB

Although variometers are not absolute, they can very accurately measure the change in the magnetic field, i.e. $\Delta B = B(t_i) - B(t_{i-1})$.

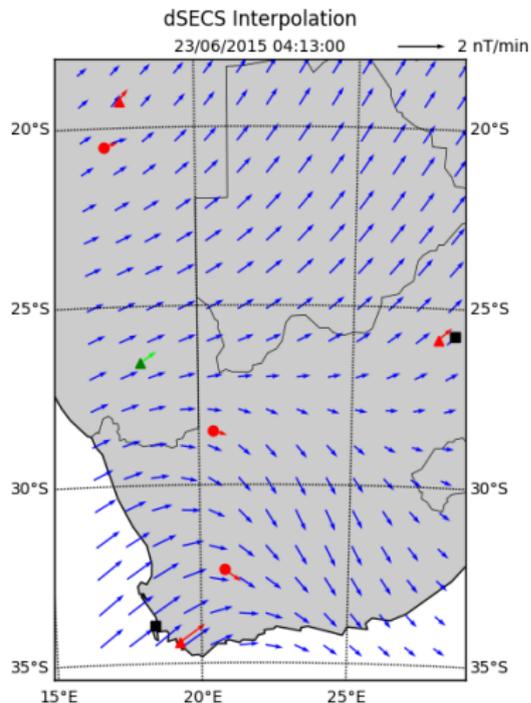
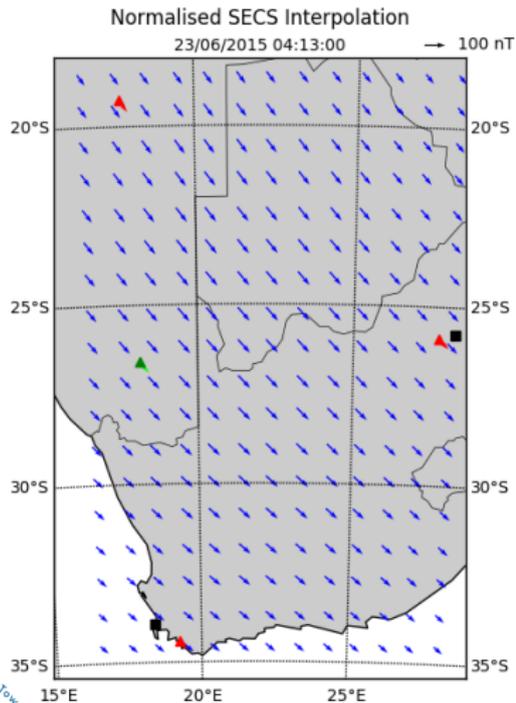
Since the SECS method is entirely linear, ΔB can be used in the same way as B (T is purely a spatial constant). The only difference is that I becomes ΔI in this scheme.

$$\begin{aligned} B(t_i) - B(t_{i-1}) &= T \cdot I(t_i) - T \cdot I(t_{i-1}) \\ &= T \cdot (I(t_i) - I(t_{i-1})) \\ \Delta B &= T \cdot \Delta I \end{aligned}$$

With more variometers than absolute magnetometers, the confidence in ΔB interpolation is much higher than that for B . ΔB is also what is typically used for GIC studies.



Merging SECS and dSECS



Merging SECS and dSECS

We can now use the greater amount of information about ΔB to improve the interpolation of B . Given a set of perturbations ϵ , the two results can be equated,

$$\Delta B_1 = (B_2 + \epsilon_2) - (B_1 + \epsilon_1)$$

...

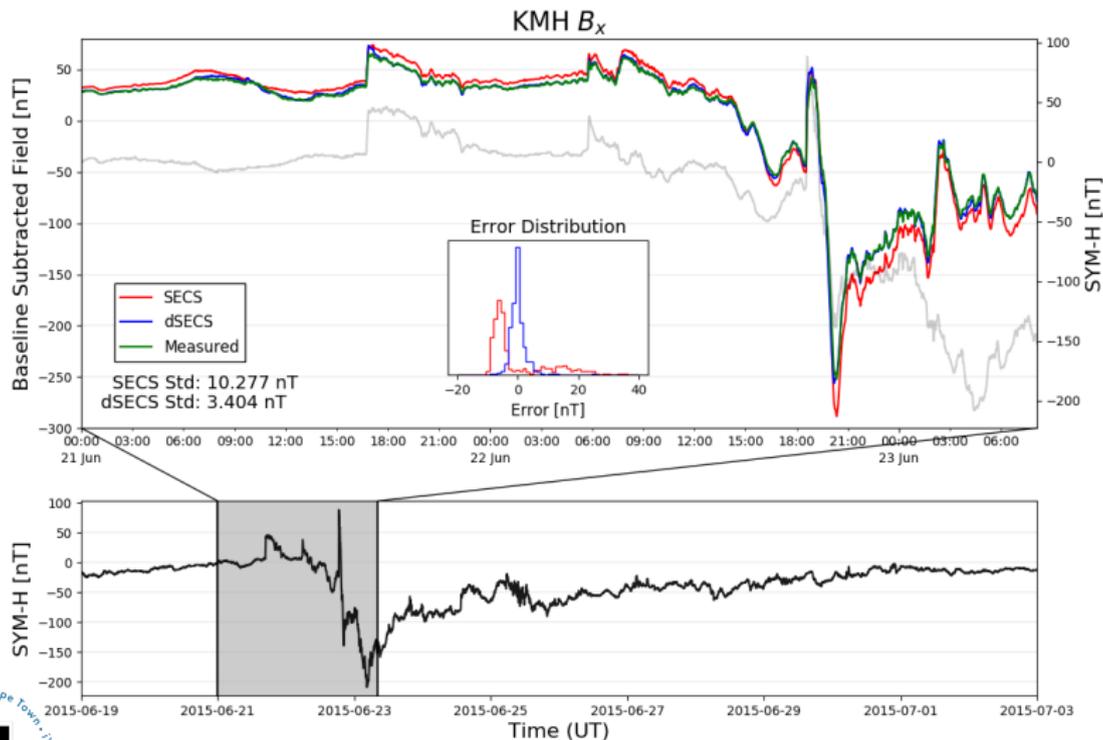
$$\Delta B_N = (B_{N+1} + \epsilon_{N+1}) - (B_N + \epsilon_N).$$

This can then be rewritten in a matrix equation of the form $A\vec{x} = \vec{b}$,

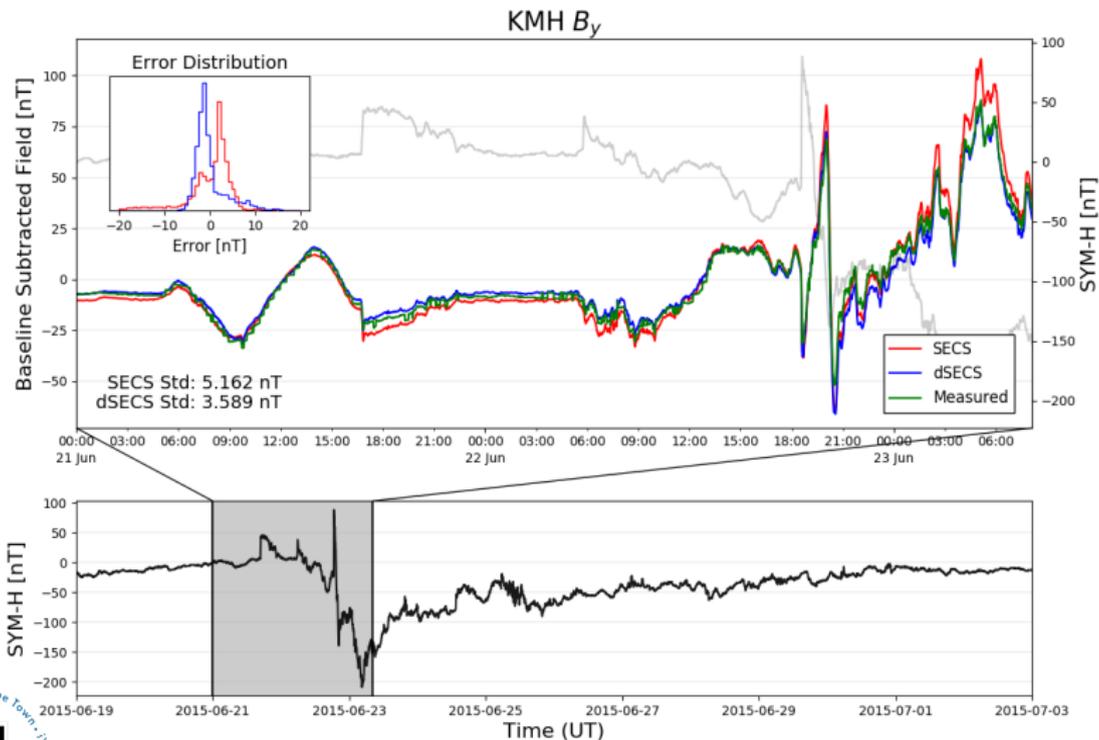
$$\begin{array}{c} \xrightarrow{N+1} \\ N \end{array} \begin{bmatrix} -1 & 1 & 0 & \dots & 0 \\ 0 & -1 & 1 & \dots & 0 \\ \vdots & \ddots & \ddots & \ddots & \vdots \\ 0 & \dots & 0 & -1 & 1 \end{bmatrix} \begin{bmatrix} \epsilon_1 \\ \epsilon_2 \\ \vdots \\ \epsilon_{N+1} \end{bmatrix} = \begin{bmatrix} \Delta B_1 + B_1 - B_2 \\ \Delta B_2 + B_2 - B_3 \\ \vdots \\ \Delta B_N + B_N - B_{N+1} \end{bmatrix}.$$



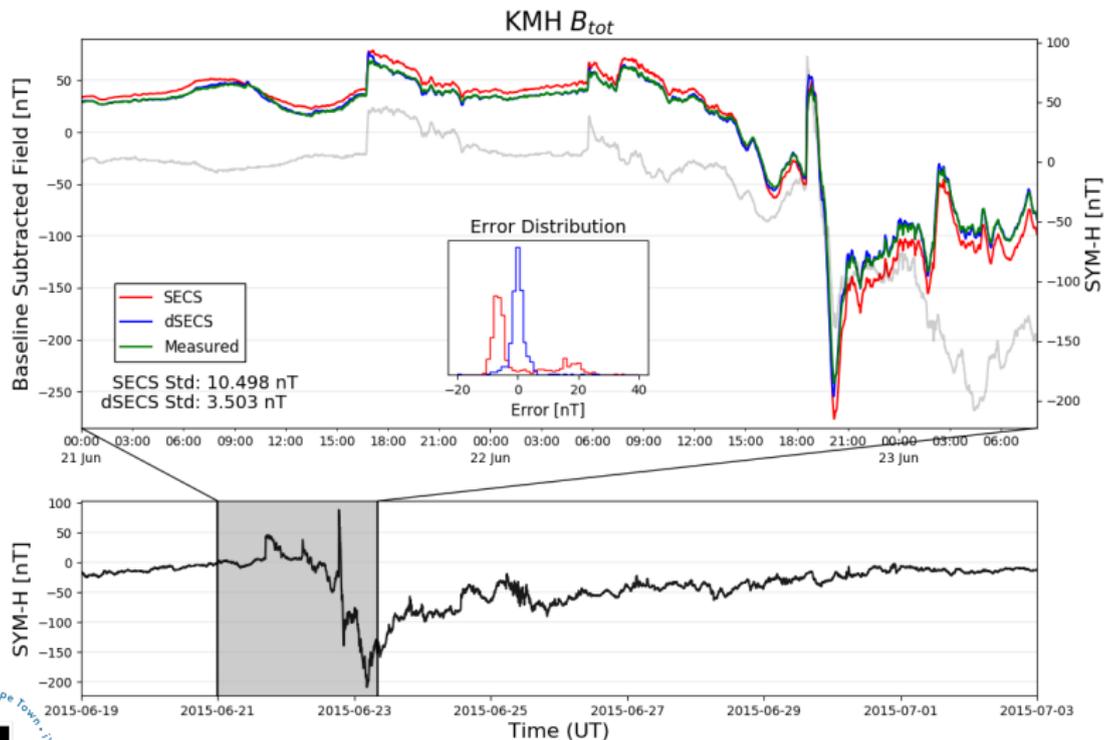
KMH B_x

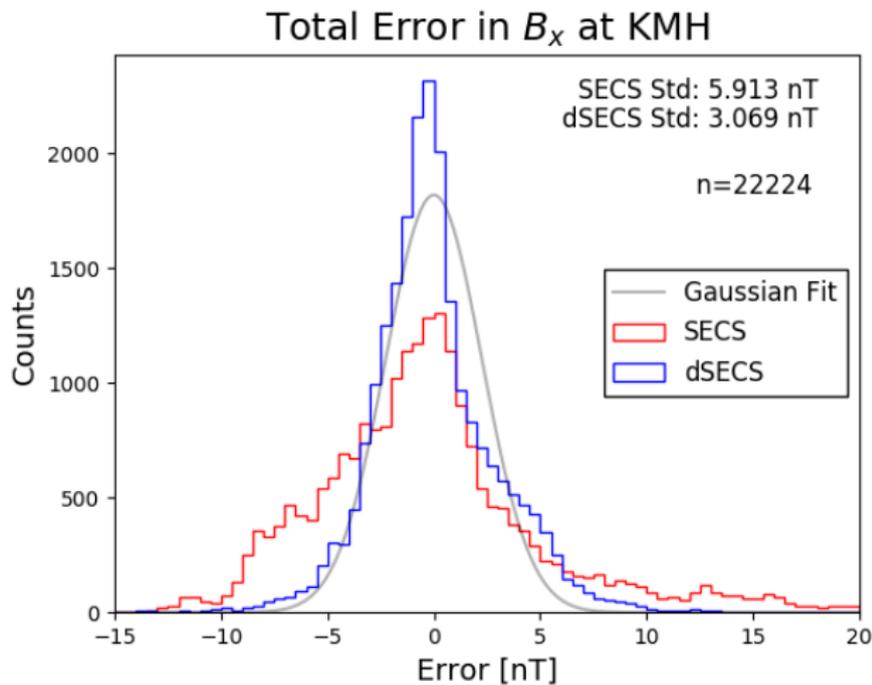


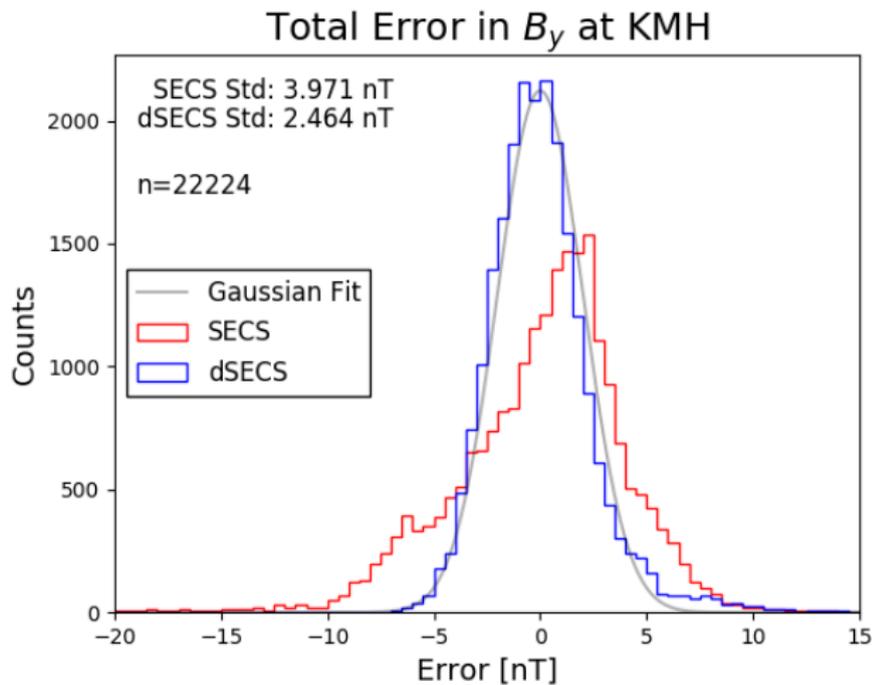
KMH B_y

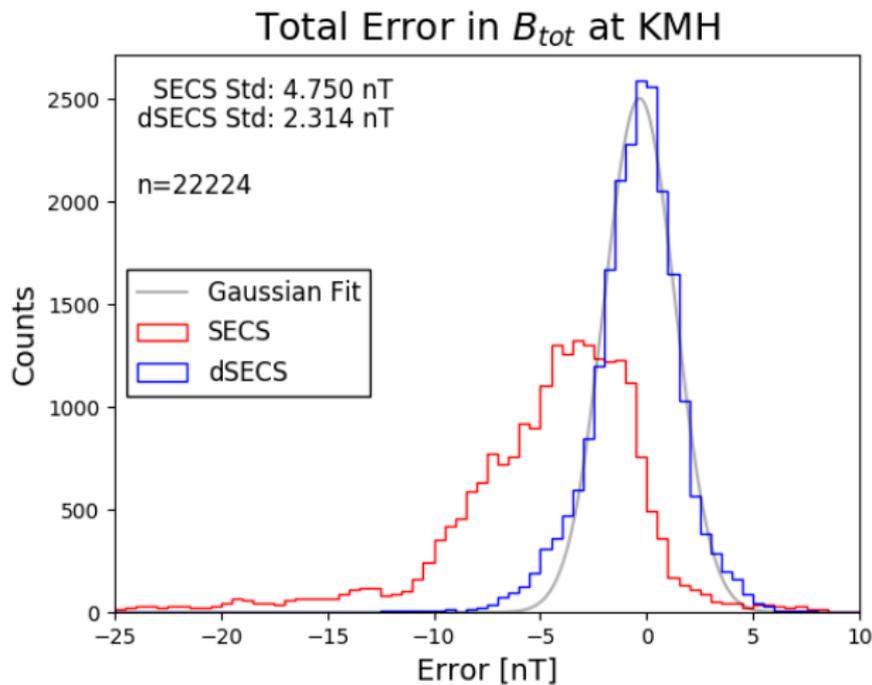


KMH B_{tot}



KMH B_x - All Storms

KMH B_y - All Storms

KMH B_{tot} - All Storms

Take Home Points...

1. Integrating variometers into magnetometer networks improve geomagnetic interpolation, especially during geomagnetic storms.
 - This improvement comes at a much reduced cost!
2. Variometers can be used alone in geoelectric field studies.
3. Systematic error is reduced, particularly in B_y component (which relates to induction effects).

'The more the merrier.'



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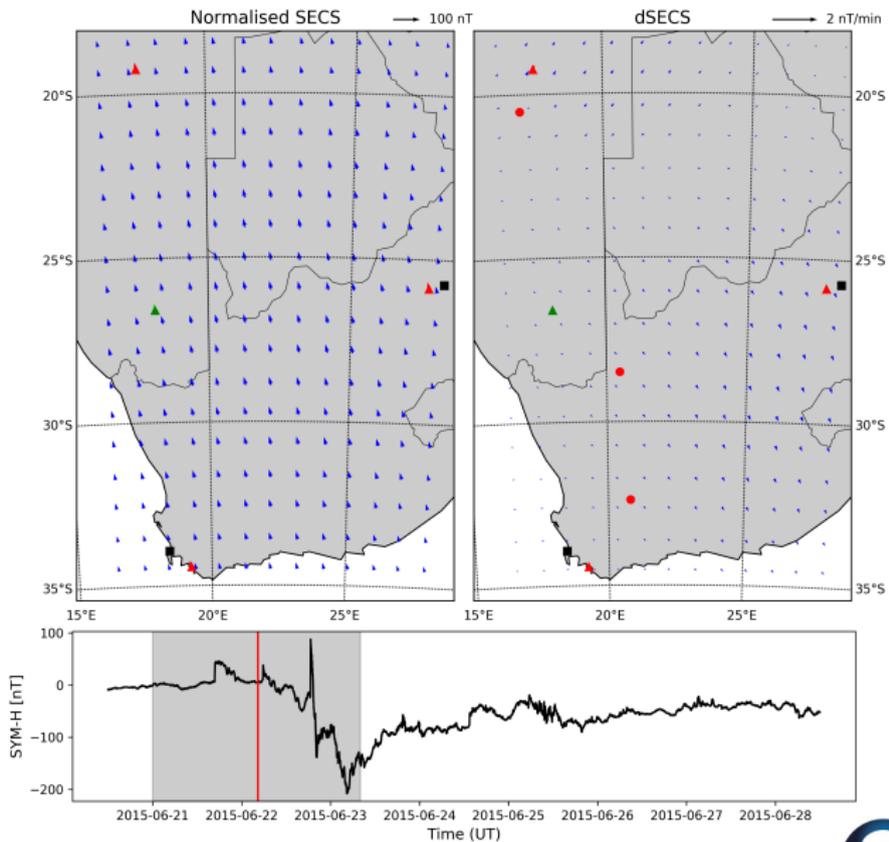
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Questions (Suggestions)?



2015-06-22 04:25:00



SECS - Background Theory

From Helmholtz's theorem, any current flowing on a surface can be broken into curl-free (allows current flow in and out of sheet) and divergence-free (allows current flow on sheet) parts. The divergence-free part is what is typically measured by ground based magnetometers.

In Earth-centred spherical coordinates (r, θ, ϕ) , the divergence-free current at a point \vec{r} on the surface R_{surf} , θ away from a pole at \vec{r}' is,

$$\vec{J}_{df,i}(\vec{r}) = \frac{I_i}{4\pi R_{surf}} \cot\left(\frac{\theta}{2}\right) \vec{e}_\phi = \frac{\iint_S \tilde{r} \cdot \vec{\nabla} \times \vec{J}(\vec{r}') dS}{4\pi R_{surf}} \cot\left(\frac{\theta}{2}\right) \vec{e}_\phi.$$

SECS - Planar Approximation

Given a current element of amplitude of I at height h in cylindrical coordinates ($r = \sqrt{x^2 + y^2}, \phi, z$), the surface current density would be,

$$\vec{J}_{df} = I/(2\pi r)\vec{e}_\phi,$$

Assuming z is downwards and a harmonic time dependence (i.e. $e^{i\omega t}$), the electric field resulting from the element would be,

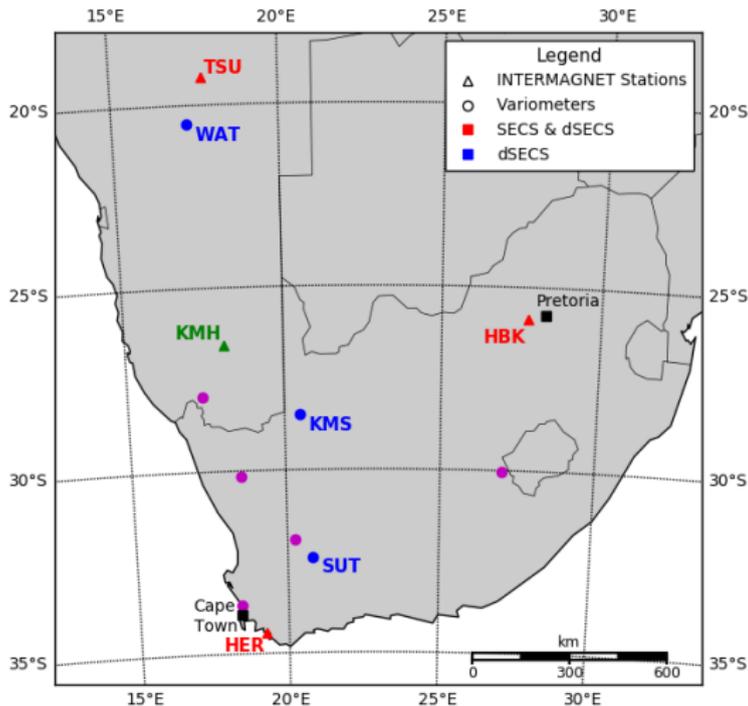
$$\vec{E} = -\frac{i\omega\mu_0 I}{4\pi} \frac{\sqrt{r^2 + h^2} - h}{r} \vec{e}_\phi.$$

The magnetic field would then be,

$$\vec{B} = \frac{\mu_0 I}{4\pi r} \left(\left(1 - \frac{h}{\sqrt{r^2 + h^2}} \right) \vec{e}_r + \left(\frac{r}{\sqrt{r^2 + h^2}} \right) \vec{e}_z \right).$$

Relevant Stations and Interpolation Grid

- Latitude range: 34.5-18.5°S
- Longitude range: 16.5-28.0°E
- Grid dimensions: 13 x 18
- Grid span: 1 214km EW and 1 781km NS
- Grid spacing: 101km and 104km
- Internal field model: Enhanced Magnetic Model (EMM2017)



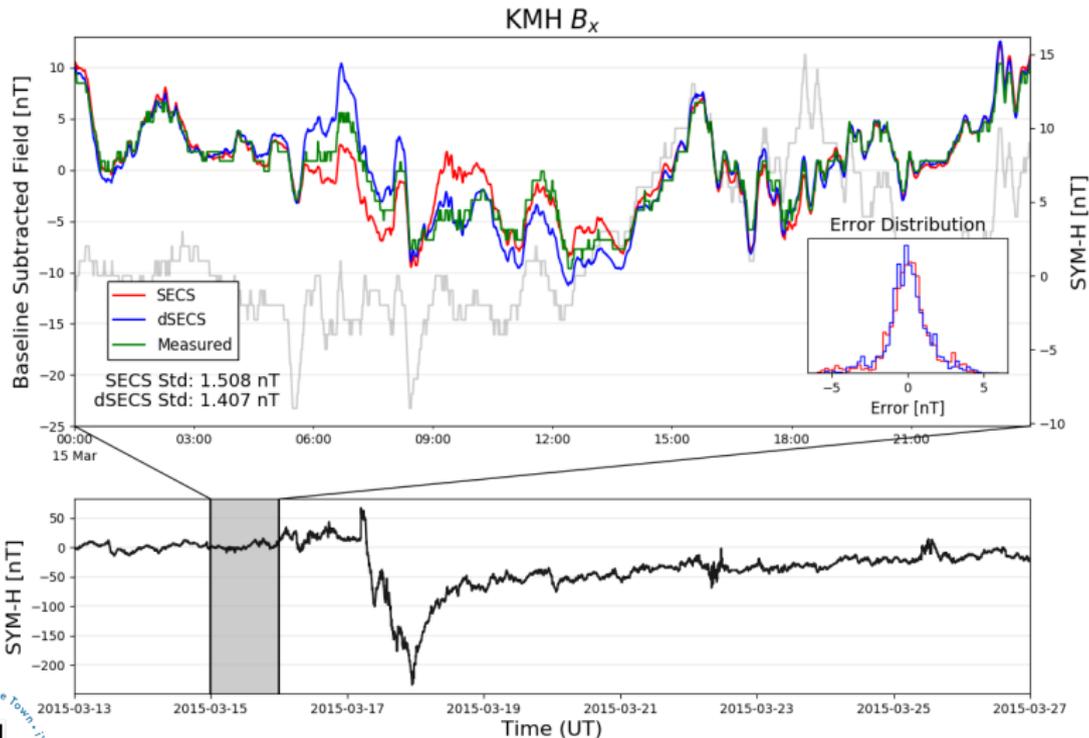
Analytical Pseudo Inverse

$$AA^{-1} = \mathbb{1}(N \times N)$$

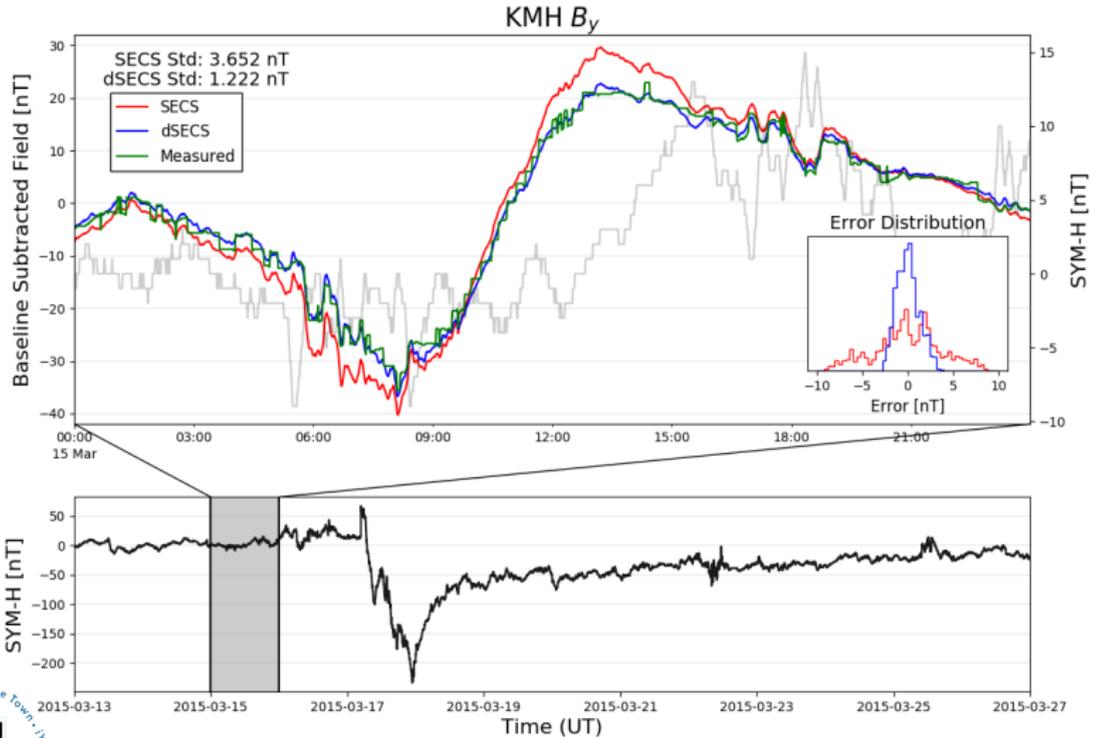
$$A^{-1}A = \begin{array}{c} \xrightarrow{N+1} \\ \left[\begin{array}{ccccc} \frac{N}{N+1} & -\frac{1}{N+1} & -\frac{1}{N+1} & \cdots & -\frac{1}{N+1} \\ -\frac{1}{N+1} & \frac{N}{N+1} & -\frac{1}{N+1} & \cdots & -\frac{1}{N+1} \\ -\frac{1}{N+1} & -\frac{1}{N+1} & \frac{N}{N+1} & \cdots & -\frac{1}{N+1} \\ \vdots & \vdots & \ddots & \ddots & \vdots \\ -\frac{1}{N+1} & -\frac{1}{N+1} & \cdots & \frac{N}{N+1} & -\frac{1}{N+1} \\ -\frac{1}{N+1} & -\frac{1}{N+1} & \cdots & -\frac{1}{N+1} & -\frac{1}{N+1} \\ -\frac{1}{N+1} & -\frac{1}{N+1} & \cdots & -\frac{1}{N+1} & \frac{N}{N+1} \end{array} \right] \downarrow N+1 \end{array}$$

$$\approx \mathbb{1}(N+1 \times N+1) \text{ for } N \gg 1$$

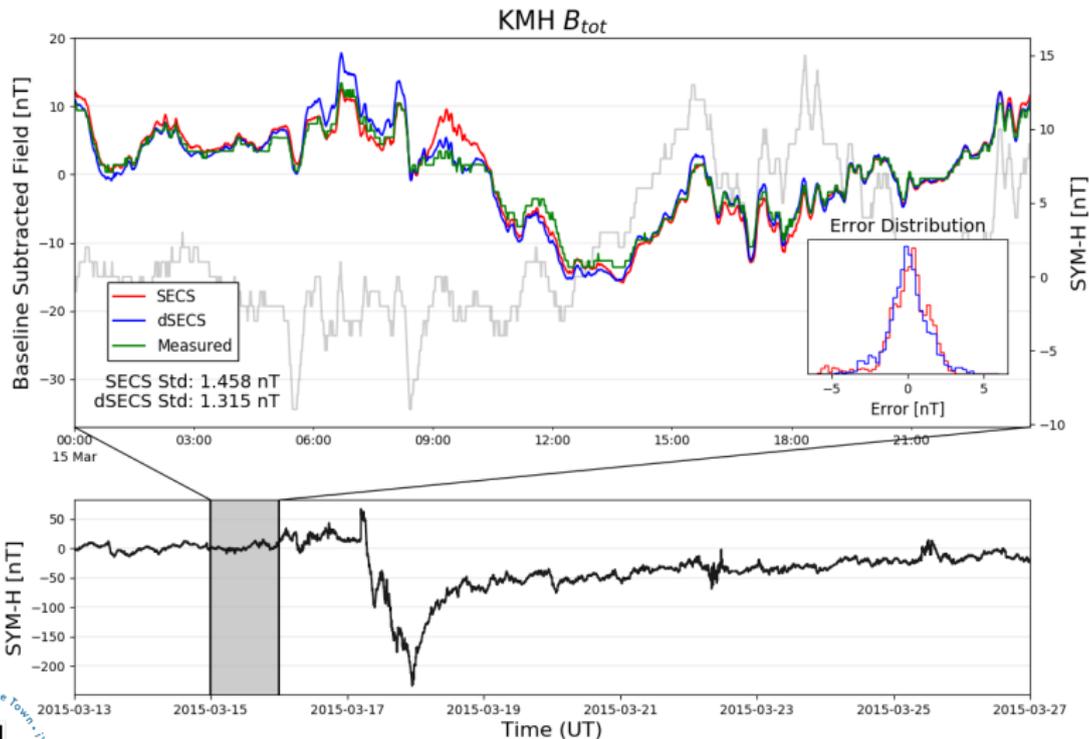
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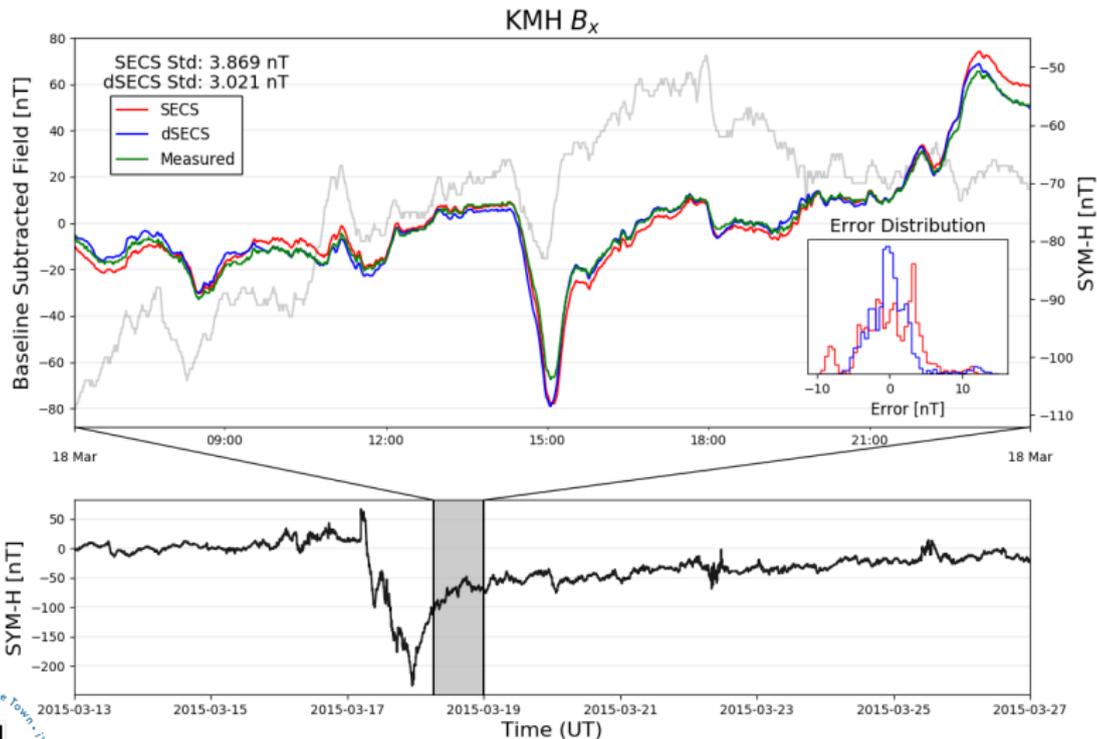
KMH B_y



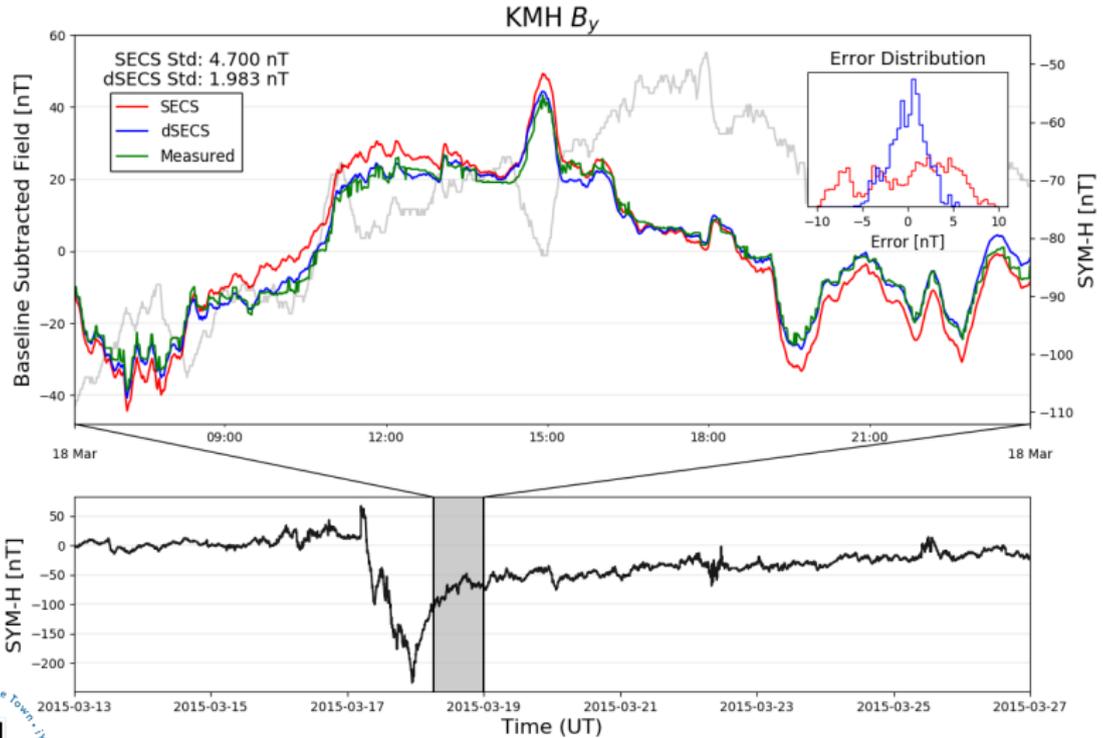
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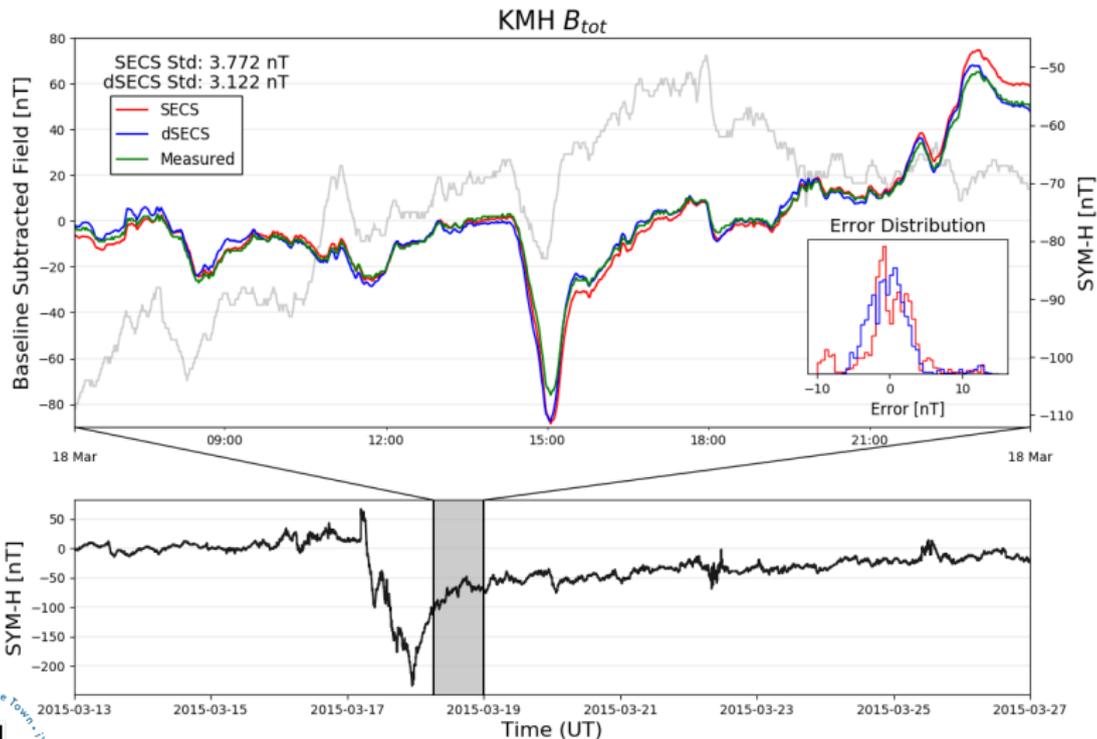
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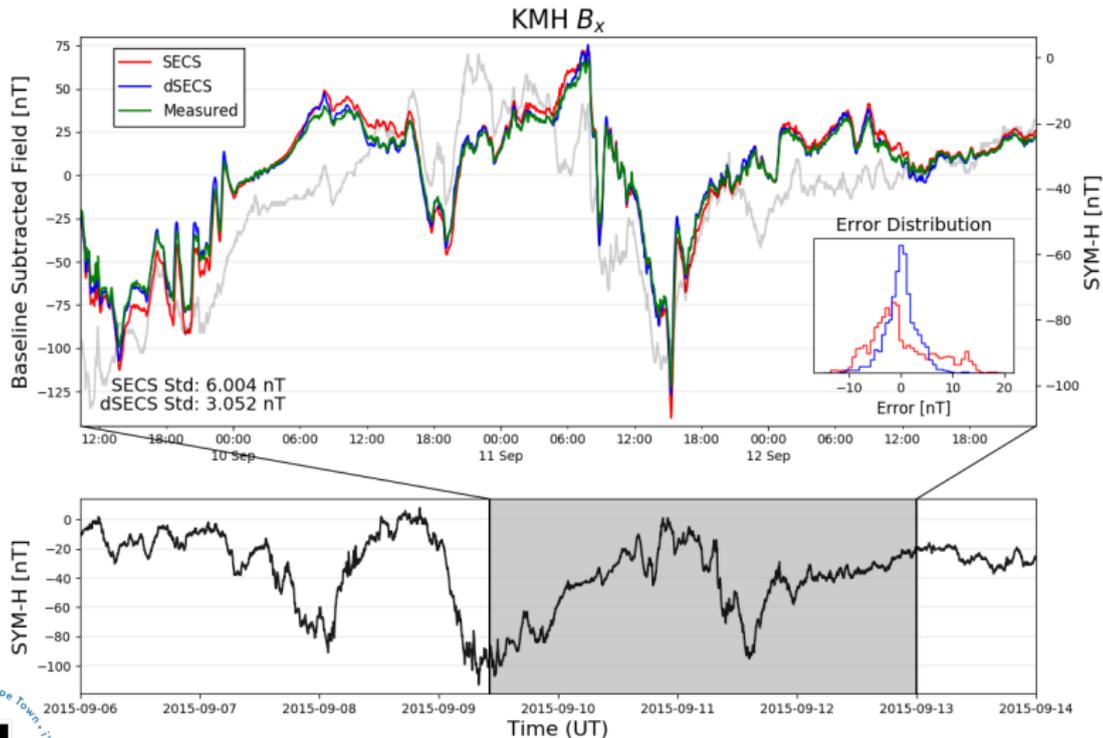
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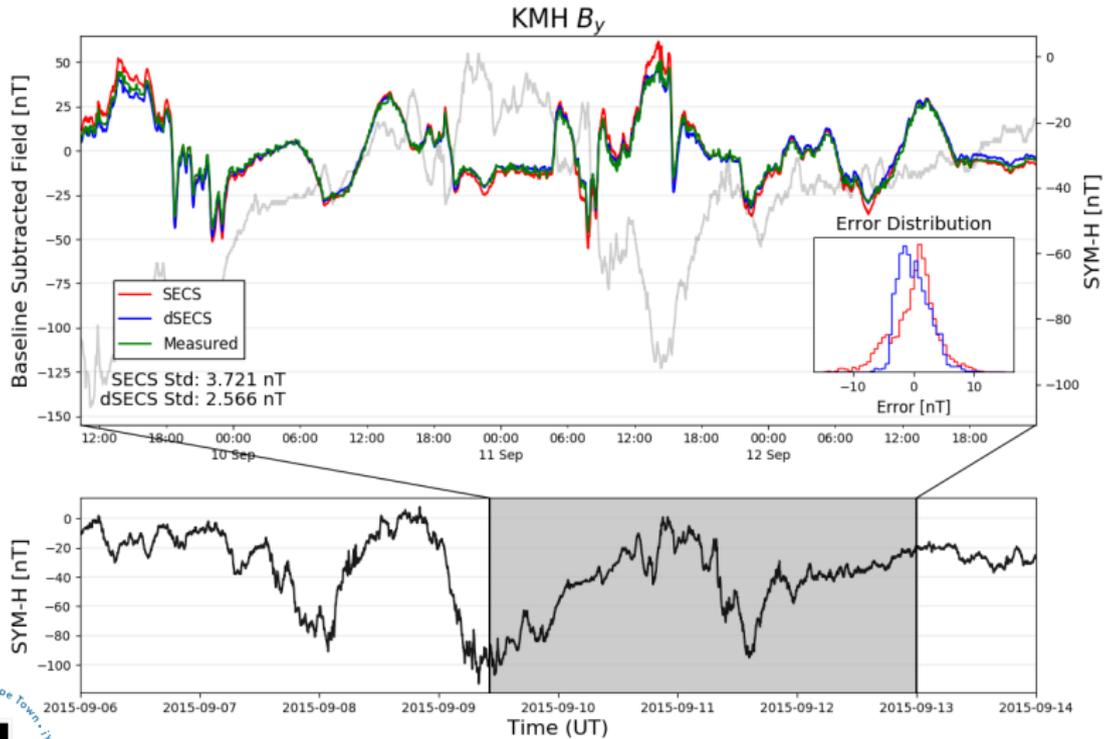
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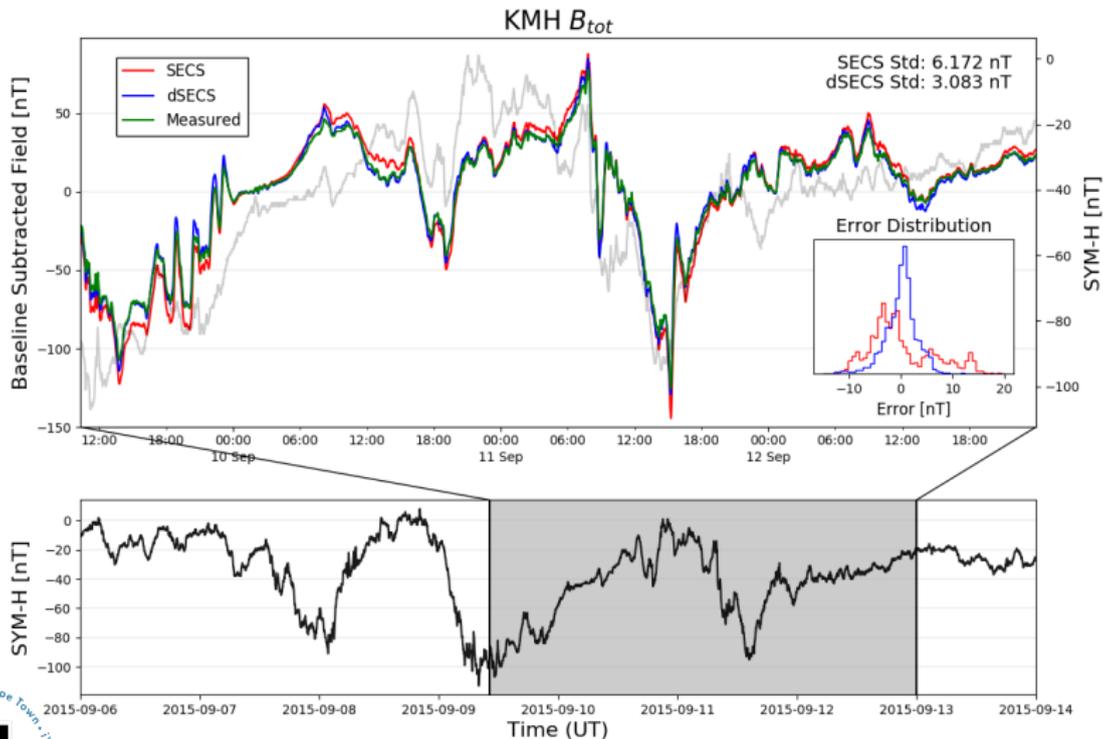
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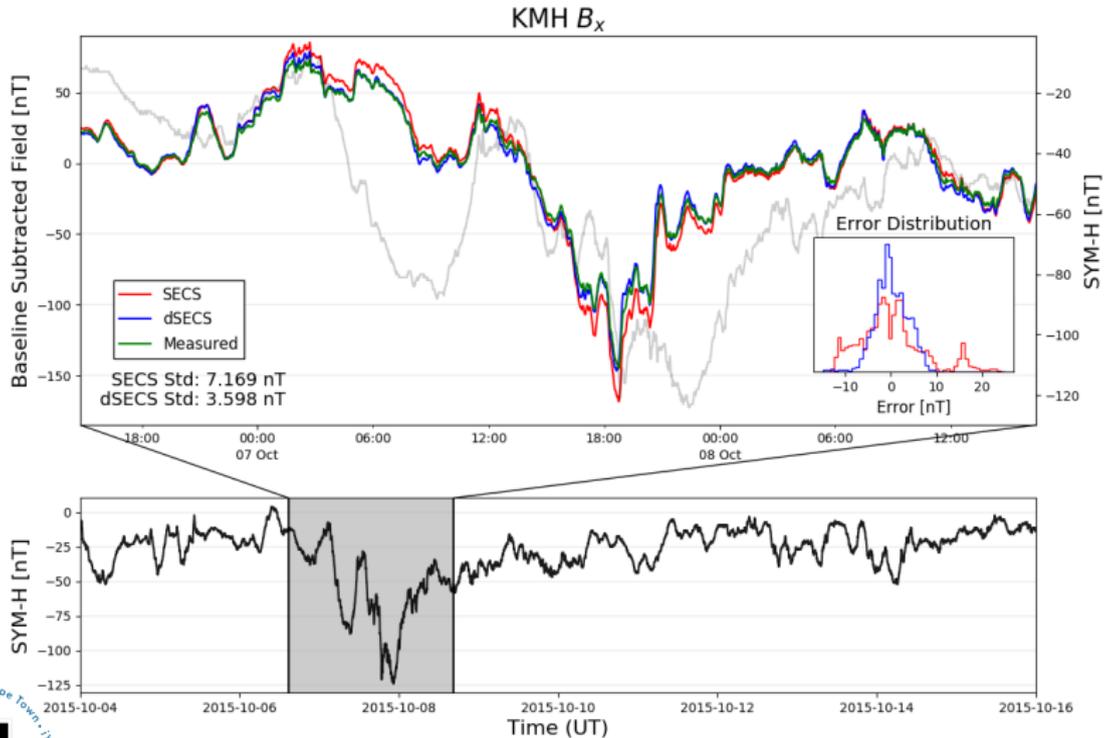
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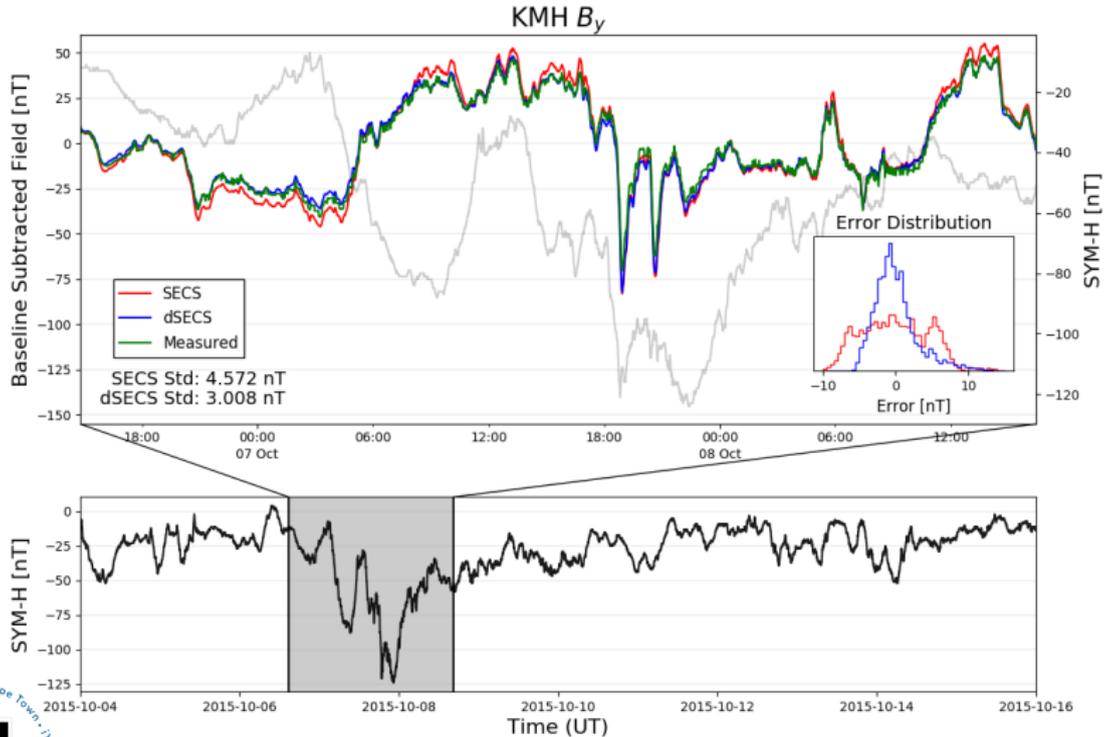
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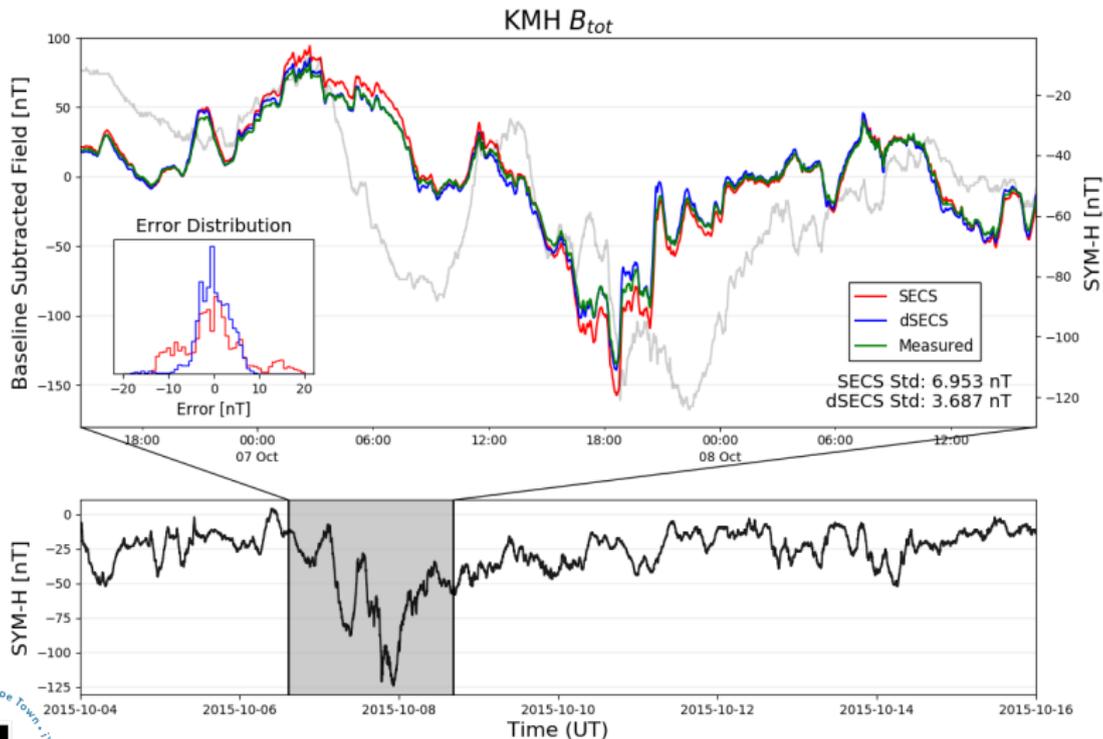
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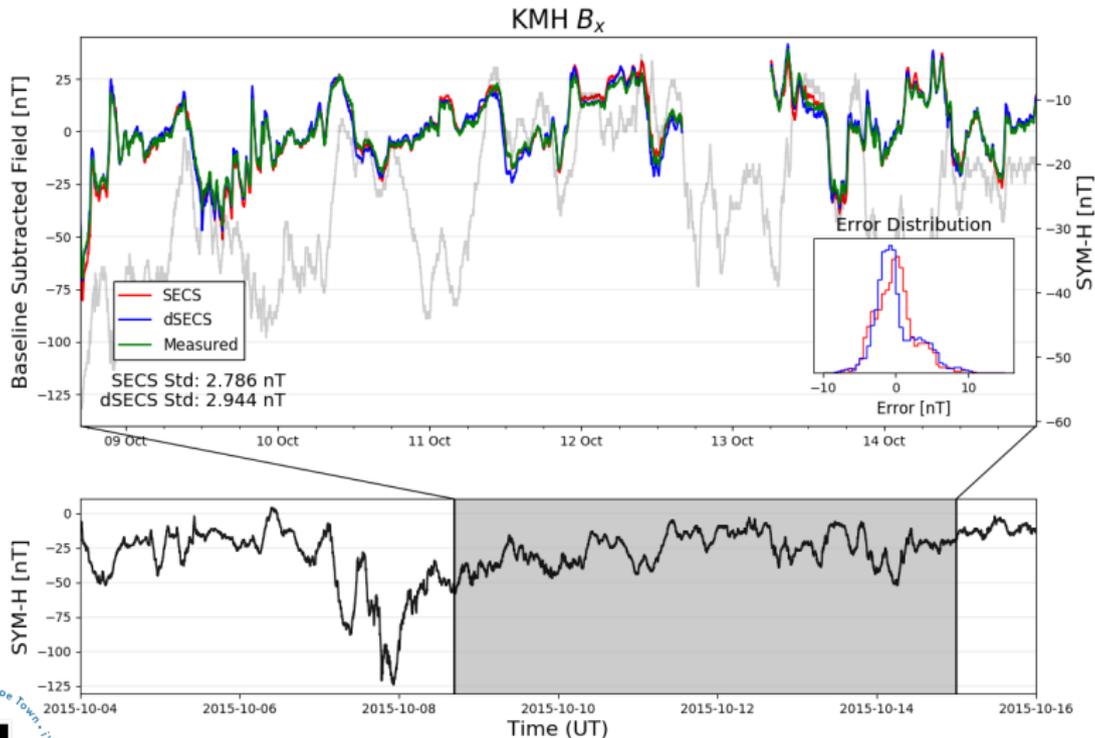
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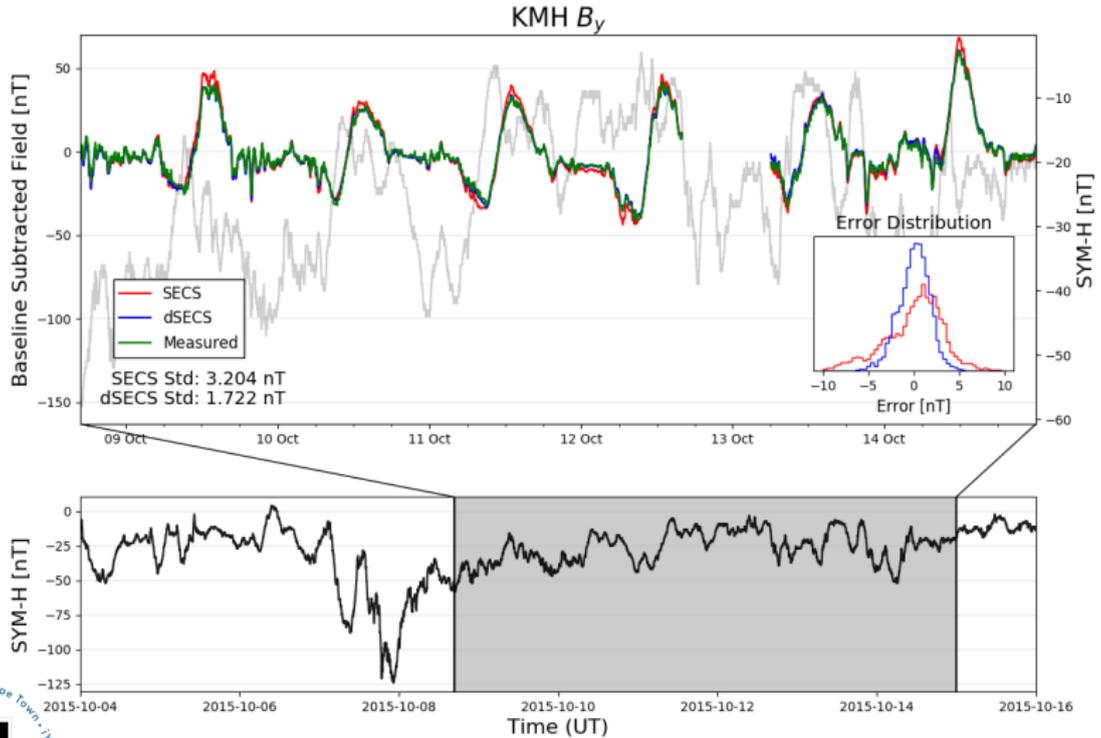
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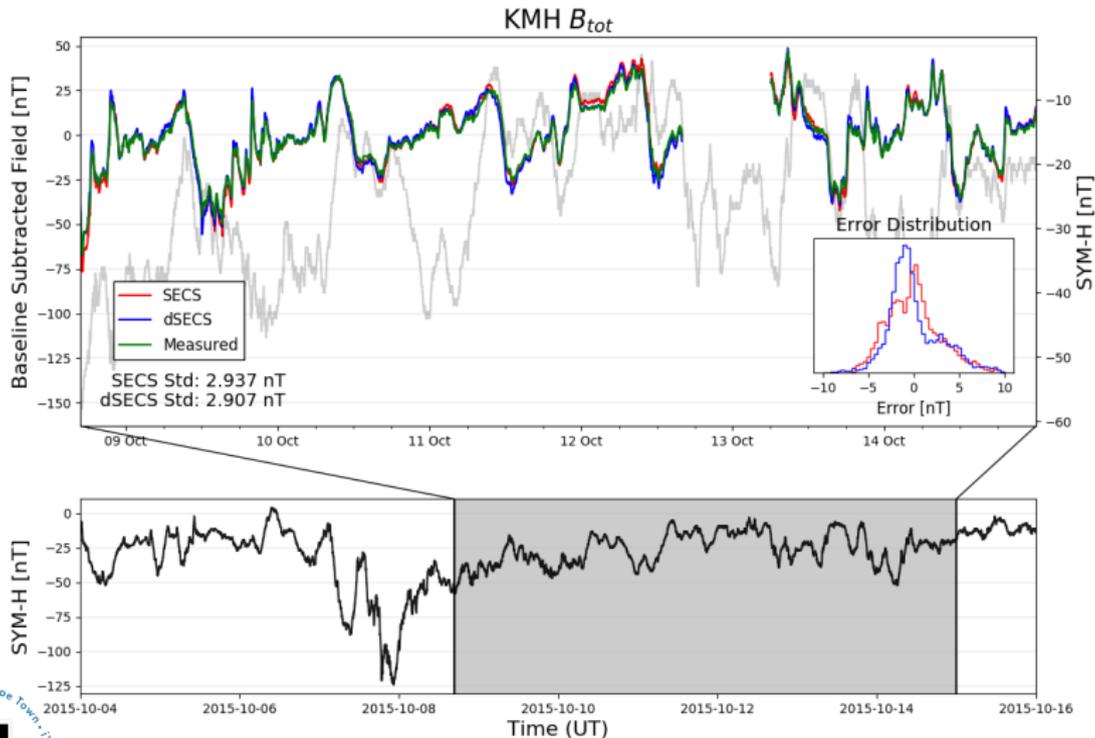
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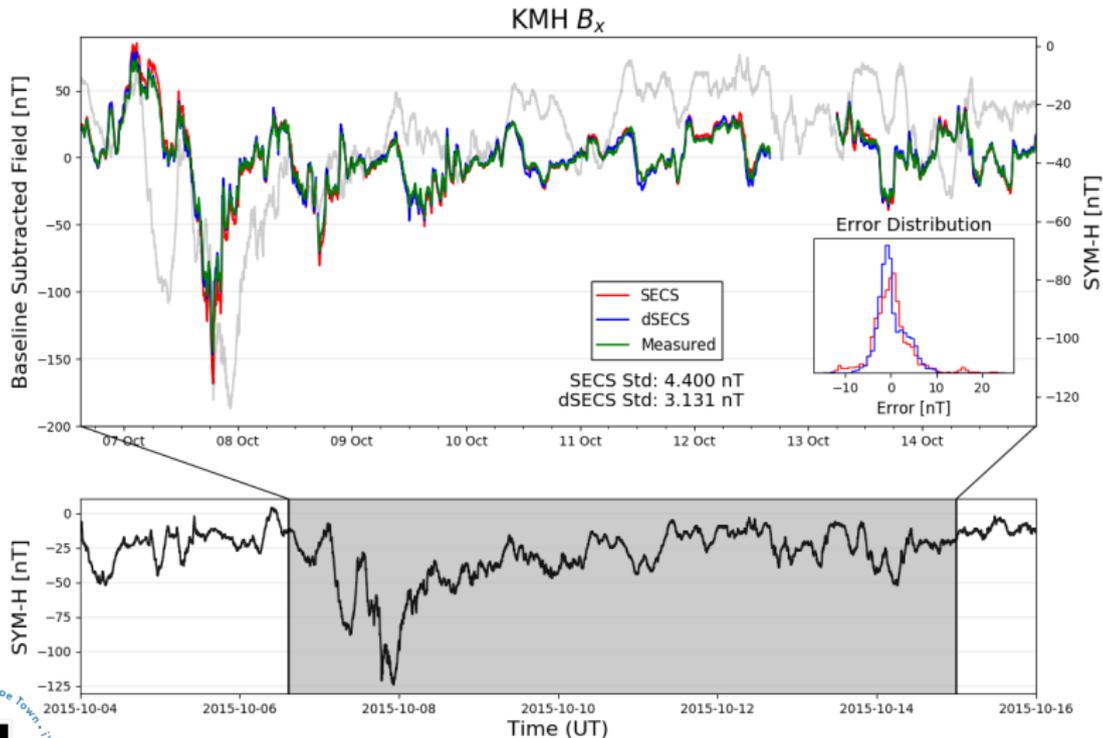
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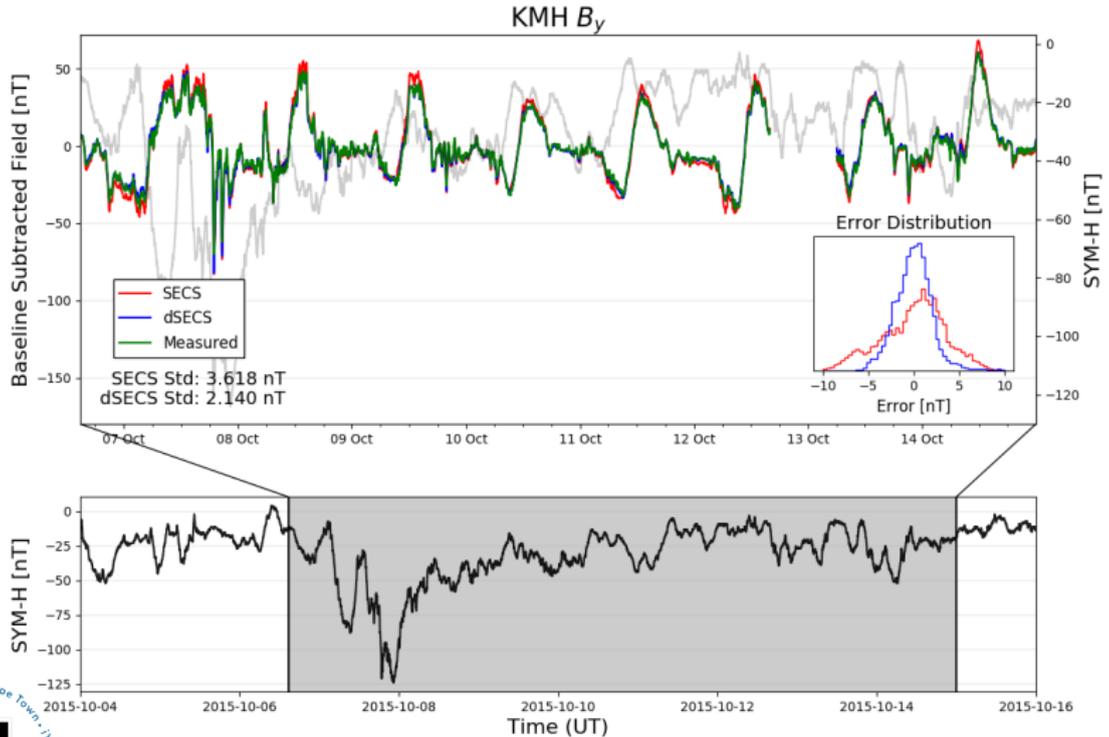
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